

# Spatio-temporal Analysis of Surface and Root Zone Soil Moisture Derived from SMAP and SMOS Microwave Satellite Data Using in-situ Measurements in Senegal

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**Abstract** Soil moisture plays a regulating role on Earth-Atmosphere flux and the water balance at the terrestrial surface. In recent years, significant progress has been made in quantifying soil moisture from satellite data. Particularly, microwave space remote sensing remains an efficient tool for estimating soil moisture at a large scale. However, the validation by in situ measurements is still needed in some areas such as the Sahel, especially for elaborated products such as soil moisture in the root zone. The present study is contributing to the validation of soil moisture products in the silvopastoral and agricultural system within the Sahelian area in Senegal. Soil moisture products derived from Soil Moisture Active and Passive (SMAP) sensors at 9 and 36 km and Soil Moisture Ocean Salinity (SMOS) at the resolution of  $0.25^\circ \times 0.25^\circ$  are compared with in-situ measurements acquired at different depths (from 0 to 100 cm) three validation sites. First, a validation of the in-situ measurements was done to visualize the performance of these sensors in situ for soil moisture recovery on the three study sites (Loumbi, Bawdi, and Niakhar). Second, an estimation of the SMAP, SMOS products of the surface area (5 cm depth) was made. Third, an estimation showed that these microwave sensors follow the soil moisture dynamics whatever the surface and climatic conditions and also despite RMSE higher or equal to the specifications,  $0.04 \text{ m}^3/\text{m}^3$ . Finally, the results from the soil moisture analysis of the root zone of the satellite and in-situ products are very important.

**Keywords:** validation, soil moisture, SMAP, SMOS, microwave remote sensing, Senegal

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## 1. Introduction

Soil moisture (SM) is an essential parameter for characterizing surface conditions relevant to the ecosystem, weather forecasting, climate prediction, agricultural productivity, flooding, and early drought warnings [1-5]. Generally expressed in gravimetric units ( $\text{g}/\text{cm}^3$ ) or volumetric units ( $\text{cm}^3/\text{cm}^3$ ), SM can be defined as the ratio of the total volume of water present in the unsaturated soil zone to the total volume of soil. According to the World Meteorological Organization, SM has been declared as an Essential Climate Variable (ECV) due to its importance in several hydrological and atmospheric processes [6].

However, SM has a strong influence on soil-vegetation-atmosphere exchanges, and its impact remains particularly high in three major regions of the world: the Great Plains of North America, India, and Sahel region, where precipitation is very strongly influenced by SM [2,4,5]. Indeed, several studies have been carried out to properly estimate this variable and its influence on soil-atmosphere exchanges [7,8,9].

However, the use of remote sensing [10-17] and land surface modeling [18-21] remain the best techniques to better understand the effects of SM between the atmosphere and terrestrial surfaces.

The potentialities of satellite measurements of SM on a near-global scale began with the Scanning Multichannel Microwave Radiometer (SMMR, C-band 6 GHz) in 1978 [22] and many algorithms based on the inversion of a

radiative transfer model have been developed for SM estimation [23,24].

Since then, several other sensors have been put into orbit such as the Advanced Microwave System Radiometer/Earth Observing System (AMSR/E-EOS) [23,25,26] in 2002. Quantification of SM with this sensor is based on X-band (10.65 GHz) brightness temperatures [10,27,28] or a combination of C-band (6.91 GHz) [29,30] and X-band brightness temperatures [31,32]. Other examples include the Microwave Imaging Radiometer with Aperture Synthesis (MIRAS), the Soil Moisture and Ocean Salinity (SMOS) satellite which provides soil moisture over the terrestrial surfaces and sea salinity over the oceans [33,34], the Soil Moisture Active and Passive (SMAP) which also provides SM and freezes/thaw values across the globe [35,36]. The Advanced SCATterometer (ASCAT) onboard the MetOP weather satellite also provides information on the degree of saturation (relative humidity) for the upper soil layer (up to 5 cm depth) [10,37,38,39].

In addition, global SM data can also be obtained from reanalysis products [40,41,42], and also from operational numerical weather prediction systems [43]. Some of these model-based products assimilate surface observations to improve the quality of SM estimates. For example, the SM-DAS-2 product assimilates ASCAT surface SM retrievals and screen-level air temperature and humidity measurements [44].

In this study, SM products from the SMAP Level 3 (L3), version 4 and 7, and Level 4 version 5 sensors and

SMOS version 650 are used. These obtained data include surface soil moisture and root zone SM estimations.

## 2. Data and Methods

### 2.1. Study Sites

Geographically, the three areas (Figure 1) are located in the Sahelian zone of Senegal, more precisely in the pastoral zone of the Ferlo for the Loumbi (15°22' N and 15°01' W) and Bawdi sites (16°02' N and 14°94' W) and in the groundnut basin at Niakhar (14°49' N and 16°46' W) for the third site. The climate of these sites is characterized by two distinct seasons: (i) a short rainy season from July to October and (ii) a long dry season from November to June. From 1950 to 1984, the total annual rainfall gradually decreased, according to [45,45,46,47] leading to a long and dry period [45,46].

The Ferlo, a sub-basin of the Senegal River, is a vast expanse of shrubby steppes and typically Sahelian savannahs south of the Senegal River valley [45]. This area, which straddles five administrative regions (Diourbel, Louga, Matam, Saint-Louis, and Tambacounda), extends approximately between latitudes 16°15' and 14°30' North and longitudes 12°50' and 16°01' West [46]. Its surface area varies according to the authors from 56,269 km<sup>2</sup>, i.e., approximately 28% of the national territory, to 60,000 km<sup>2</sup> or 70,000 km<sup>2</sup>, which makes it one of the largest eco-geographical zones in Senegal [45,46].

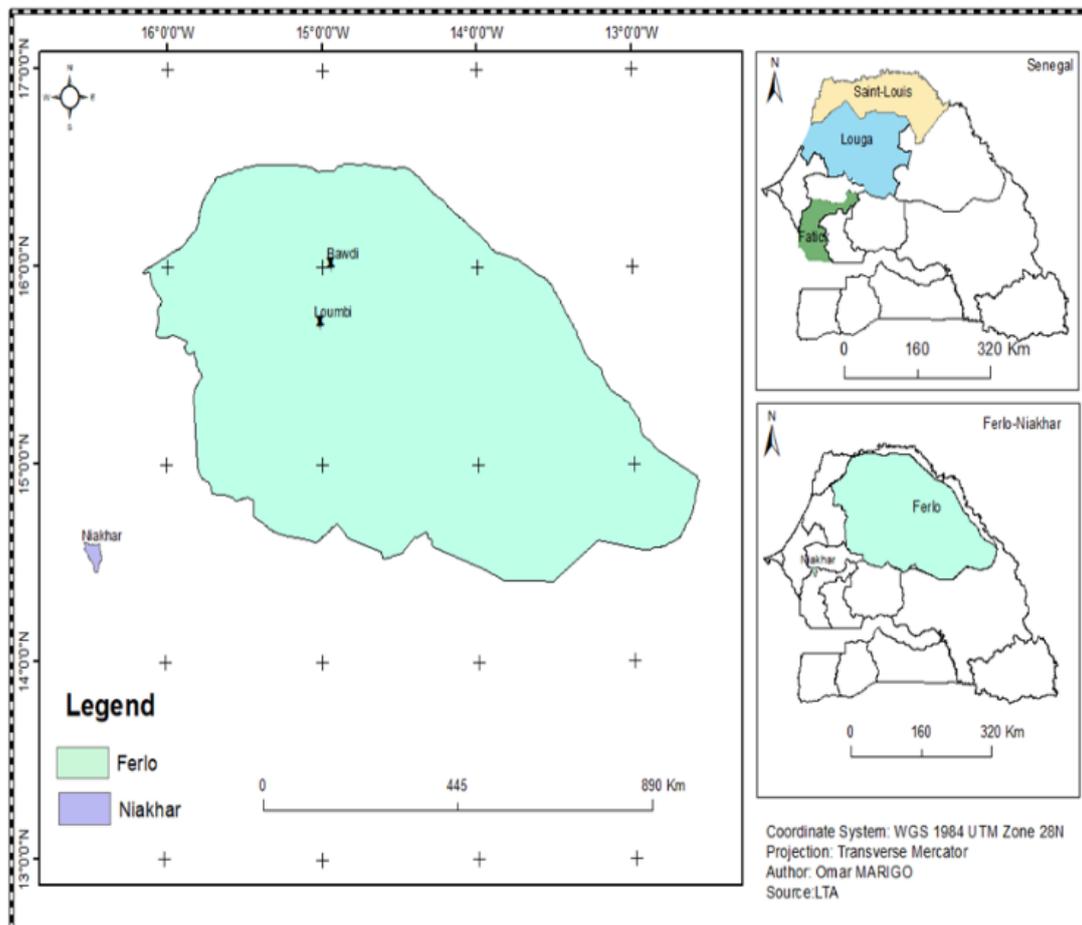


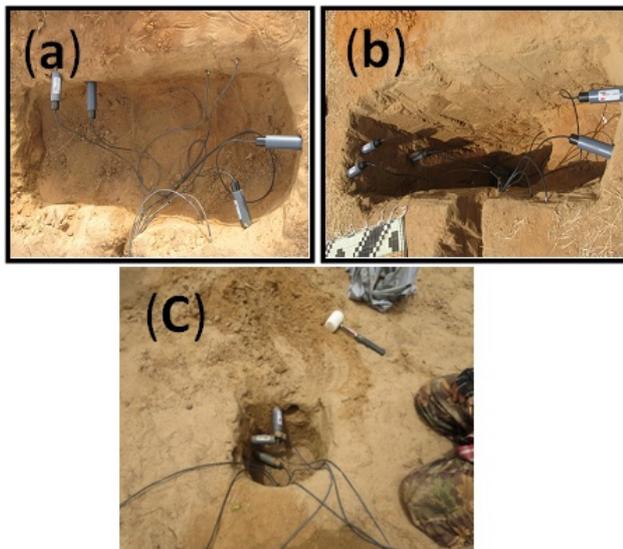
Figure 1. Location of the three sites in the study area (Ferlo: Bawdi and Loumbi; Niakhar)

The soil map of the Ferlo is dominated by sandy textured soil [45,46]. The reliefs are separated by longitudinal depressions with greyish sandy-clay soil, locally chalky and hydromorphic soil with temporary waterlogging, as well as sub-arid brown-red soil. Annual rainfall in the Ferlo is between 200 and 500 mm, associated with high interannual variability (coefficient of variation: 30 to 40%), low relative air humidity in the dry season, and high during the rainy season (annual average of around 35%), high air temperatures (annual average of 29°C), high evaporation (1800 to 2200 mm/year) and precariousness of soil water reserves [46,48,49].

The Niakhar site is located in the administrative region of Fatick, down the groundnut basin of Senegal. It is characterized by a tropical Sudanian climate with rainfall varying between 600 and 900 mm per year on average. Minimum air temperatures range from 21°C to 22°C, while maximum air temperatures range from 35°C to just over 36°C. The soils of the region belong to main 4 types: tropical ferruginous soils (Dior and Deck), hydromorphic soils of the valleys, holomorphic soils (saline or "tanne" soils), and mangrove soils found in the islands and estuaries. The region's landscapes are highly artificial, with a predominance of rain-fed farming areas (84%) (CSE: Centre de Suivi Ecologique, 2015).

## 2.2. Description of Used Data

The data for this study is composed of field SM measurements and SM products derived from SMAP (Soil Moisture Active and Passive), SMOS (Soil Moisture Ocean and Salinity) sensors, and precipitation data from CHIRPS (Climate Hazards Group Infrared Precipitation with Stations) sensors.



**Figure 2.** Installation of soil moisture and temperature sensors at: (a)-Bawdi, (b)-Loumbi and (c)-Niakhar

### 2.2.1. In Situ Measurements

The in-situ data are obtained from SM ML2 Theta Probe (TP) sensors buried underground at different depths (Figure 2, Table 1). These probes are sensitive to the variation of the soil dielectric constant which is then

converted into SM via a calibration relationship, the latter being a function of the soil type [50]. The sensors automatically record SM and temperature measurements with a fifteen minutes step. For the three study sites with sandy soil (> 90% sand), we applied the calibration provided by the manufacturer:

$$SM_1 = \frac{100}{a_1} * 4.70 * \left(\frac{s_1}{1000}\right)^3 - 6.40 * \left(\frac{s_1}{1000}\right)^2 + 6.40 * \left(\frac{s_1}{1000}\right) - a_0 + 1.00 \quad (1)$$

With  $SM_1$ : calibrated soil moisture;  $s_1$ = raw data;  $a_0$ =1.6 and  $a_1$ =8.4 ( $a_0$  and  $a_1$  are the calibration constants).

In this study, we set up ML2 probes at different depths at the three study sites (Table 1). Subsequently, the 5 cm and 10 cm depth probes were used in the Ferlo (Bawdi and Loumbi) and Niakhar respectively for the evaluation and analysis of SMAP surface soil moisture products with resolutions of 36 km (SPL3SMP), 9 km (SPL3SMP\_E), and 9 km (SPL4SMGP\_5). All the other depths are used in the estimation and validation of the root zone SM at the Loumbi and Niakhar sites.

**Table 1.** Site names, number of implemented SM sensors, and sites coordinates

| Sites   | Number of probes/Soil Moisture | Lat/Long     | Period of analysis |
|---------|--------------------------------|--------------|--------------------|
| Bawdi   | 2                              | 16°02/ 14°94 | 2015- 2019         |
| Loumbi  | 6                              | 15°72/ 15°01 | 2015- 2019         |
| Niakhar | 3                              | 14°49/ 16°45 | 2018- 2019         |

### 2.2.2. Satellites Data

Table 2 summarizes the characteristics of the satellite products used, their launch dates, their spatial-temporal resolutions, and their accessibility.

#### 2.2.2.1. SMAP Products

Launched in 2015, with a temporal resolution of 2-3 days, the SMAP radiometer provides daily global soil moisture products that are generated by the SMAP Science Data Processing System (SDS) at JPL (Jet Propulsion Laboratory) [12,51,52,53].

The estimation of soil moisture from surface Brightness Temperature (BT) measurements is based on the inversion of the radiative transfer equation, commonly known as the tau-omega model of soil moisture by passive microwaves [54].

The SDS provides three types of soil moisture products at different spatial resolutions with an accuracy of 0.04 m<sup>3</sup>/m<sup>3</sup>:

- The level 3 products (SPL3SMP) with a spatial resolution of 36 km x 36 km. They are derived from the SMAP Level-1C TB (L1C\_TB) product which contains the calibrated, geo-located, and time-ordered L1B\_TB brightness temperatures that have been resampled to the fixed 36 km grid EASE2.0 (Equal-Area Scalable Earth Grid, Version 2.0) [54,55]. To obtain an accurate estimate of soil moisture, several auxiliary data sets are used. These auxiliary data sets include surface temperature, the

optical thickness of vegetation, simple scattering albedo of vegetation, surface roughness information, land use type, soil texture, and data indicators for identification of land, water, precipitation, urban areas, mountainous terrain, and dense vegetation [54,56,57].

- Enhanced level 3 products (SPL3SMP\_E) with a spatial resolution of 9 km x 9 km [54]. This product is a daily SMAP Level-2 (L2) soil moisture composite that is derived from interpolated SMAP Level-1C (L1C) brightness temperatures. Backus-Gilbert’s optimal interpolation techniques are used to extract information from the SMAP antenna temperatures and convert them to brightness temperatures, which are displayed on the 9 km x 9 km Elementary Earth Grid, version 2.0 (EASE-Grid 2.0) in a global cylindrical projection [54].
- Level 4 Surface and Root-Zone Soil Moisture data (L4\_SM) which provide a surface (0-5cm) and root-zone (0-100cm) soil moisture estimates at 9 km x 9 km resolution [58,59]. Data generated using a terrestrial data assimilation system that combines the advantages of spatial measurements of L-band brightness temperatures, precipitation observations, and land surface modeling. The L4\_SM data are therefore generated and distributed on the global grid, cylindrical at 9 km resolution (Equal-Area Scalable Earth, version 2 EASE\_2) [59].

**2.2.2.2. SMOS Products**

The SMOS (Soil Moisture and Ocean Salinity) satellite is an Earth exploration mission of the European Space Agency (ESA) in collaboration with the French National Space Agency (CNES) and the Spanish Centre for the Development of Industrial Technology (CDTI) [33,60,61]. Launched in November 2009, it was the first radiometer operating at L-band (1.4 -2 GHz) [62,63,64,65]. Its main objective is to provide global surface soil moisture at a depth of 0-5 cm, with a spatial resolution of 0.25° x 0.25°, a revisit time of fewer than 3 days, and an accuracy of 0.04 m<sup>3</sup>/m<sup>3</sup> [62].

In the present study, the SM data sets available at the CATDS (Centre de Traitement des Données SMOS) are used. The CATDS consists of an operational processing center (C-PDC at IFREMER) and a SM C-EC at the CESBIO (Centre d’Etudes Spatiales de la Biosphère), Toulouse [66]. The objective of CATDS is to produce and distribute level 3 and 4 products, to propose the improvement of level 2 soil moisture algorithms using temporal information (level 3), and to produce higher level datasets from the combination of SMOS data with

physical models or other remote sensing products (level 4) to estimate root zone SM. The data used is available in the link In Table 2.

**2.2.2.3. Rainfall Products: CHIRPS**

CHIRPS (Climate Hazards Group Infrared Precipitation with Stations) is a near-global rainfall data set of over 40 years [67]. CHIRPS integrates satellite images with a spatial resolution of 0.05° x 0.05° with in situ station data to create gridded precipitation time series for trend analysis and seasonal drought monitoring.

The CHIRPS algorithm includes the following steps [67,68,69]: (a) derive TIR (Thermal Infrared) precipitation estimates from near-global geostationary satellite observations, which are generated using local regressions between the multi-satellite pentad tropical precipitation mission and the Cold Cloud Duration (CCD); (b) converting the TIR to percentage anomaly and multiplying by CHPclim (Climate Hazards Group Precipitation Climatology), producing unbiased precipitation fields. Step (a) yields the CHIRPS product, which is a time series back to 1981 at a spatial resolution of 0.05° latitude/longitude. Table 3 summarizes the annual rainfall totals from 2015 to 2019 at the different study sites.

**2.3. Methods**

SM products from SMAP and SMOS microwave sensors are evaluated at different temporal and spatial scales; in situ measurements served as a reference and are used in the evaluation of soil moisture from microwave products and precipitation data are used to delineate the existing seasons in Senegal. The analysis is organized in three steps:

- A study of the seasonality of in situ SM observations overall sites from 2017 to 2019 with measurements recorded at 5, 10, 20, 40, and 100 cm depth. For the three validation sites, the 5 cm depth probe measurements are compared with the satellite sensor products for the study of surface moisture and the 10, 20, and 100 cm depths for the SM of the root zone.
- For the three study sites, the 5 cm depth probe measurements are compared with the satellite surface SM products (SPL3SMP, SPL3SMP\_E, SPL4SMGP\_5, SMOS) over five years and a statistical study is carried out to assess their performances.
- SMAP L4\_SM products and SMOS L4\_RZSM products are evaluated at the Loumbi and Niakhar sites taking into account the integrated soil moisture of the root zone (0 - 100 cm).

**Table 2. The set of satellite sensors used and their characteristics**

| Satellites | Period        | Spatial Resolution | Time Resolution | Accessibility   |
|------------|---------------|--------------------|-----------------|---|
| SMAP 36km  | 2015- Current | 36 x 36 km         | 2 to 3 days     | <a href="https://lpdaacsvc.cr.usgs.gov/appears/">https://lpdaacsvc.cr.usgs.gov/appears/</a>   |
| SMAP 9km   | 2015- Current | 9 x 9 km           | 2 to 3 days     | <a href="https://lpdaacsvc.cr.usgs.gov/appears/">https://lpdaacsvc.cr.usgs.gov/appears/</a>   |
| SMOS       | 2009- Current | 0.25° x 0.25°      | 2 to 3 days     | <a href="https://www.catds.fr/Products/Available-products-from-CPDC">https://www.catds.fr/Products/Available-products-from-CPDC</a><br>Google Earth Engine Platform |
| CHIRPS     | 1981- Current | 0.05° x 0.05°      | 1 day           |   |

**Table 3. Annual rainfall totals from 2015 to 2019 in mm**

| Annual Rainfall (mm) |       |        |         |
|----------------------|-------|--------|---------|
| Period               | Bawdi | Loumbi | Niakhar |
| 2015                 | 265   | 311    | ----    |
| 2016                 | 268   | ----   | ----    |
| 2017                 | 204   | 252    | ----    |
| 2018                 | 250   | 289    | 415     |
| 2019                 | 162   | 197    | 448     |

The estimation of SM products from satellite sensors was carried out by taking into account only the morning orbits (descending 6 / Am). Indeed, many studies have shown that measurements made in the morning are of better quality than those acquired in the afternoon [70,71,72]. In the morning, the soil is close to water balance and the SM is very representative of the water content of the deeper layers [71,73].

To evaluate and compare the performance of the products used four metrics-widely used by any scientists the study of SM [20,74,75] are deployed: the Pearson's correlation coefficient (R), the Root Mean Squared Error (RMSE, m<sup>3</sup>/m<sup>3</sup>), the Mean Absolute Error (MAE, m<sup>3</sup>/m<sup>3</sup>) and the bias (m<sup>3</sup>/m<sup>3</sup>) defined as below:

$$R = 1 - \frac{\sum_{i=1}^n (\theta_i - \hat{\theta}_i)^2}{\sum_{i=1}^n (\theta_i - \bar{\theta})^2} \quad (2)$$

$$Bias = \frac{\sum_{i=1}^n (\hat{\theta}_i - \theta_i)}{N} \quad (3)$$

$$RMSE = \sqrt{\frac{\sum_{i=1}^n (\hat{\theta}_i - \theta_i)^2}{N}} \quad (4)$$

$$MAE = \frac{\sum_{i=1}^n (|\hat{\theta}_i - \theta_i|)}{N} \quad (5)$$

Where  $\hat{\theta}_i$  is the satellite soil moisture measurement;  $\bar{\theta}_i$ , the average of satellite soil moisture measurements;  $\theta_i$  the in-situ soil moisture measurement and N, the total number of samples taken into consideration.

### 3. Results and Discussion

#### 3.1. In Situ Temporal Variations of SM

Figure 3 and Figure 4 show the temporal evolution of soil moisture measured at different depths profiles (5, 10, 20, 40, and 100 cm) and rainfall at the three study sites (Bawdi, Loumbi, and Niakhar) from 2017 to 2019. The lack of data from the Niakhar site is due to technical problems. An increase in SM can be noted for all probes from the first rains except for the 40 and 100 cm depth probes in Loumbi and the 20 and 40 cm depths in Niakhar.

For years of study, there is an increase in moisture in all probes from the first rains except for the 40 and 100 cm depths in Loumbi (Figure 3 (a1 and a2)) and the 20 and 40

cm depths in Niakhar (Figure 4(c1 and c2)). At the Bawdi site (Figure 3 (b1 and b2)), all probes reacted as the first rains fell, and this continued for the entire study period. These results demonstrate the performance of the ML2 probes in measuring soil moisture values at different depths. The 40 cm probe at Loumbi records an increase in humidity from August onwards, unlike the 100 cm probe, which reacts in mid-September.

This is consistent with the infiltration time of soil moisture during the rainy season. Thus, the recorded SM follows the seasonality since in the dry season they do not show significant increases. On the other hand, in the rainy season, we see distinct responses of the in-situ sensors to the acquisition of SM. We also note that in the dry season, the moisture values at depths of 40-100 cm at the Loumbi site are much higher than the moisture values measured at depths of 5 and 10 cm.

At Niakhar, three probes at depths of 10, 20, and 40 cm are installed. The measurements recorded by these probes in the rainy season show that they follow the dynamics of the existing seasons in the area. The probe installed at 40 cm records soil moisture measurements just after the 10 and 20 cm probes; this indicates faster infiltration. This may be the result of a soil difference or the plowing effect that facilitates the infiltration of rainwater. In fact, in the Niakhar site, which is an agricultural area, the soil is plowed at the first rains come, which makes it more permeable to water.

#### 3.2. Comparative Analysis of SMAP and SMOS SM Products

In both Bawdi and Loumbi sites, there are two probes of 5 cm depth, only the probe installed at a depth of 5 cm with the smallest standard deviation is used by the site to evaluate SMAP and SMOS satellite products and the 10 cm probe is used in Niakhar.

Figure 5 shows the seasonal variation of the different SM products (SMAP at 9-36 km resolution and SMOS at 0.25° x 0.25°) in comparison with in situ data. The statistical scores are presented in Table 4. Comparing annually the soil moisture results, the highest correlation coefficient is found in the Bawdi site in 2016 where R is equal to 0.96 between the SPL3SMP sensor and field measurements. In Loumbi, the highest correlation score was recorded between the SPL3SMP\_E sensor and field measurements, where R was 0.94 in 2018. 0.89 is the highest correlation coefficient recorded between the SPL3SMP\_E sensor and field measurements in Niakhar in 2019.

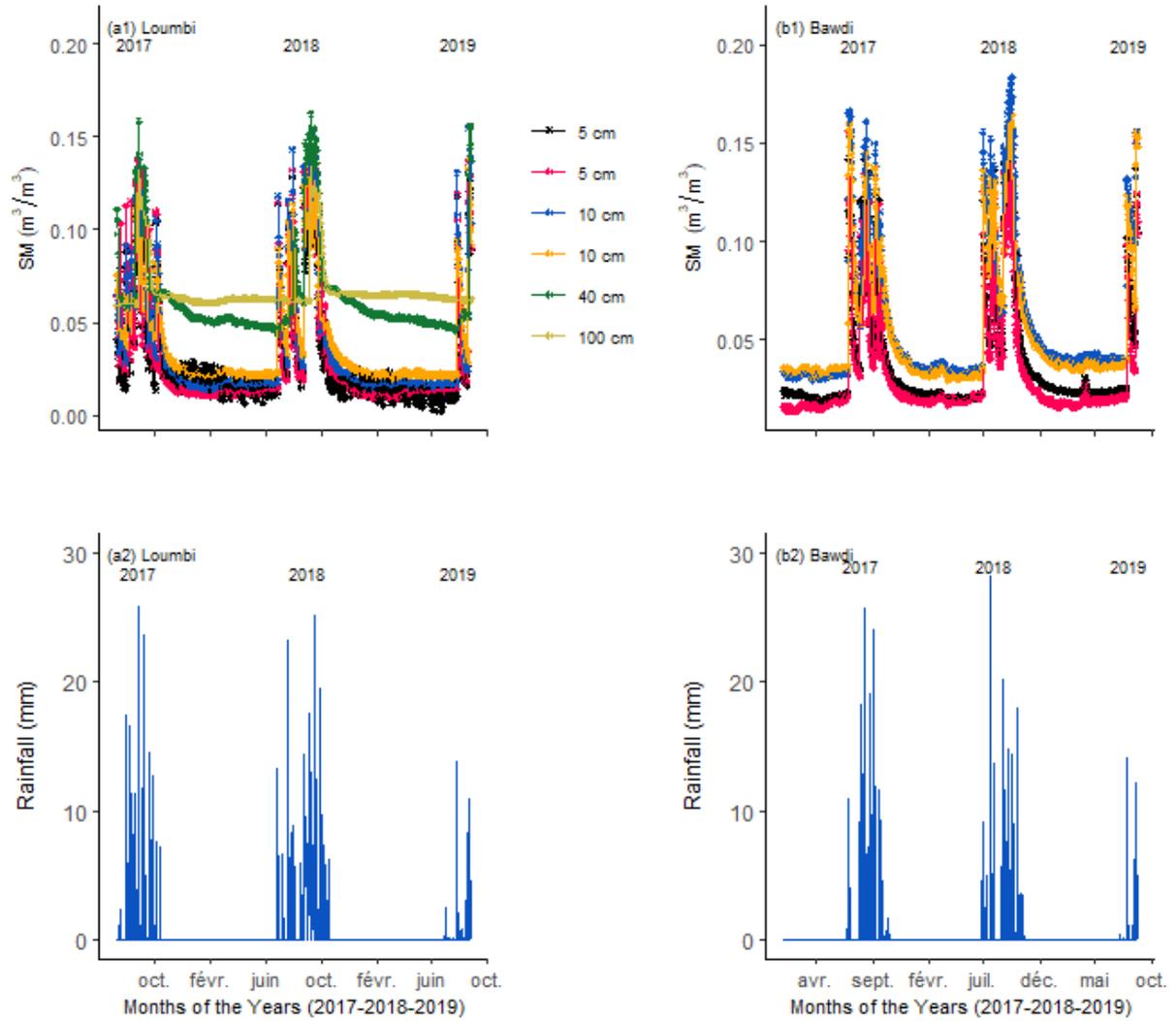


Figure 3. Time series of in situ SM measured at different depths at (a) Bawdi, (b) Loumbi sites over the 2017 – 2019 period: Pastorale zone

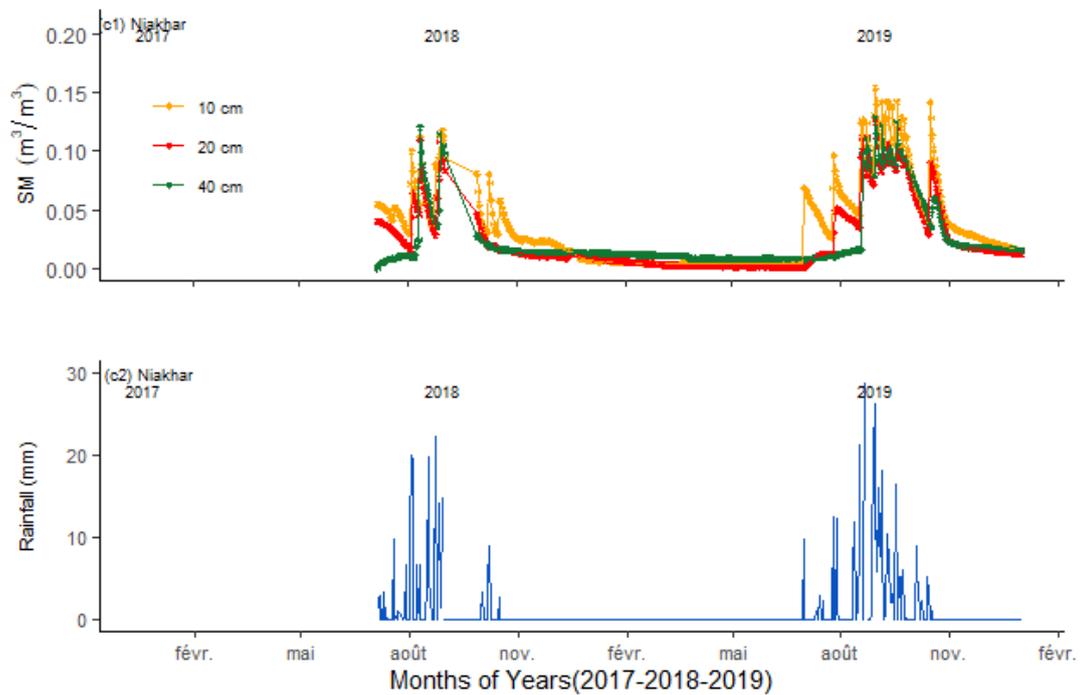
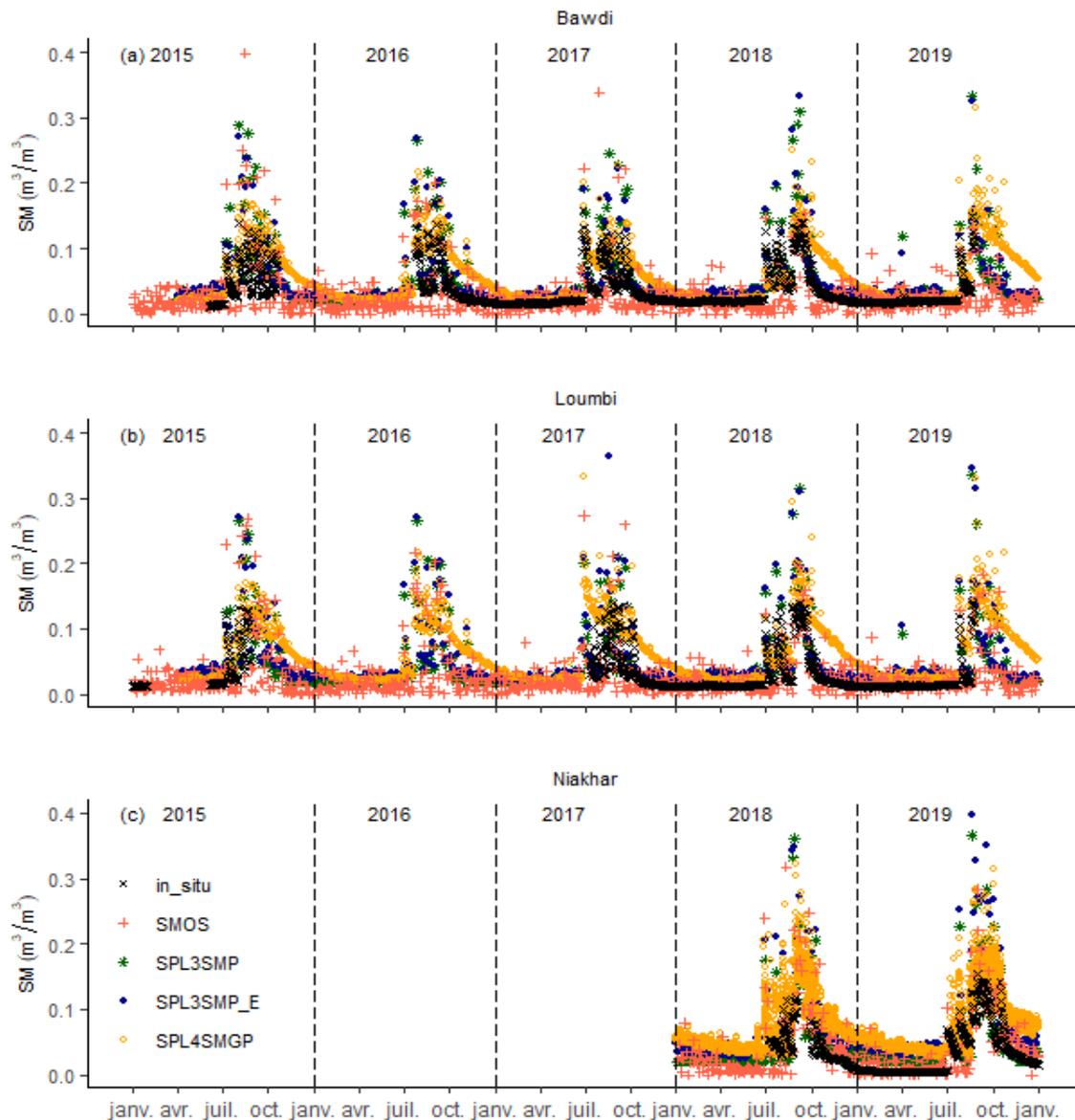


Figure 4. Time series of in situ SM measured at different depths at (c) Niakhar site over the 2017 – 2019 period: agricultural zone



**Figure 5.** Temporal variation of SMAP, SMOS, and in-situ SM data at (-a) Bawdi, (b) Loumbi and (c) Niakhar sites

In the three sites, at the end of the rainy season, Bawdi (Figure 5a), Loumbi (Figure 5b), and Niakhar (Figure 5c), the soil moisture values of the microwave sensors decrease to their minimum which is close to  $0.02 \text{ m}^3/\text{m}^3$  and the in-situ observations decrease to  $0.01 \text{ m}^3/\text{m}^3$ .

This drop lasts from January to the end of June. In fact, the end of the dry season is accompanied by dry winds and high temperatures [56,76], which accelerate the drying out of the soil, leading to a more rapid decrease in soil moisture. The SMAP products (SPL3SMP and SPL3SMP\_E) remain quite high in the rainy season and can reach maxima of  $0.3 \text{ m}^3/\text{m}^3$  soil moisture compared to the SMAP Level 4 (SPL4SMGP\_5), SMOS, and in situ products. The observations of the SPL4SMGP product are marked by a progressive decrease in soil moisture at the end of the rainy season in contrast to the other products of the two satellites studied and the field measurements. This shows either a strong sensitivity of the SMAP products to low humidity or sensitivity to other unidentified seasonal phenomena. The SMOS measurements, on the other hand, are much more scattered, especially during the dry season.

These results corroborate studies done in another site in the Ferlo, namely Dahra [77] about 80 km south of the Loumbi and Bawdi sites. They also confirm results from other areas of the Sahel, specifically Mali, Benin, and Niger among others [16,22,25,48,78,79]. This demonstrates that these L-band satellites are sensitive to the spatial-temporal dynamics of the amount of water in the soil. This moisture is strongly governed by fluctuations in rainfall, temperature, evapotranspiration, and other intrinsic soil parameters [37,80].

In Niakhar, the values of the satellite products reflect the dynamics of two existing seasons. Compared to the sites in the Ferlo, the peak of soil moisture is reached in September. The microwave measurements are much higher than the field measurements. This may mean that the satellite products recover soil moisture data better than the in-situ measurements.

The correlation between the in-situ measurements and the satellite products allows a better assessment of the accuracy of the soil moisture of the SMAP and SMOS sensors. Initially, only data at 5 cm depth were considered in this section (Figure 6, Figure 7, and Figure 8).

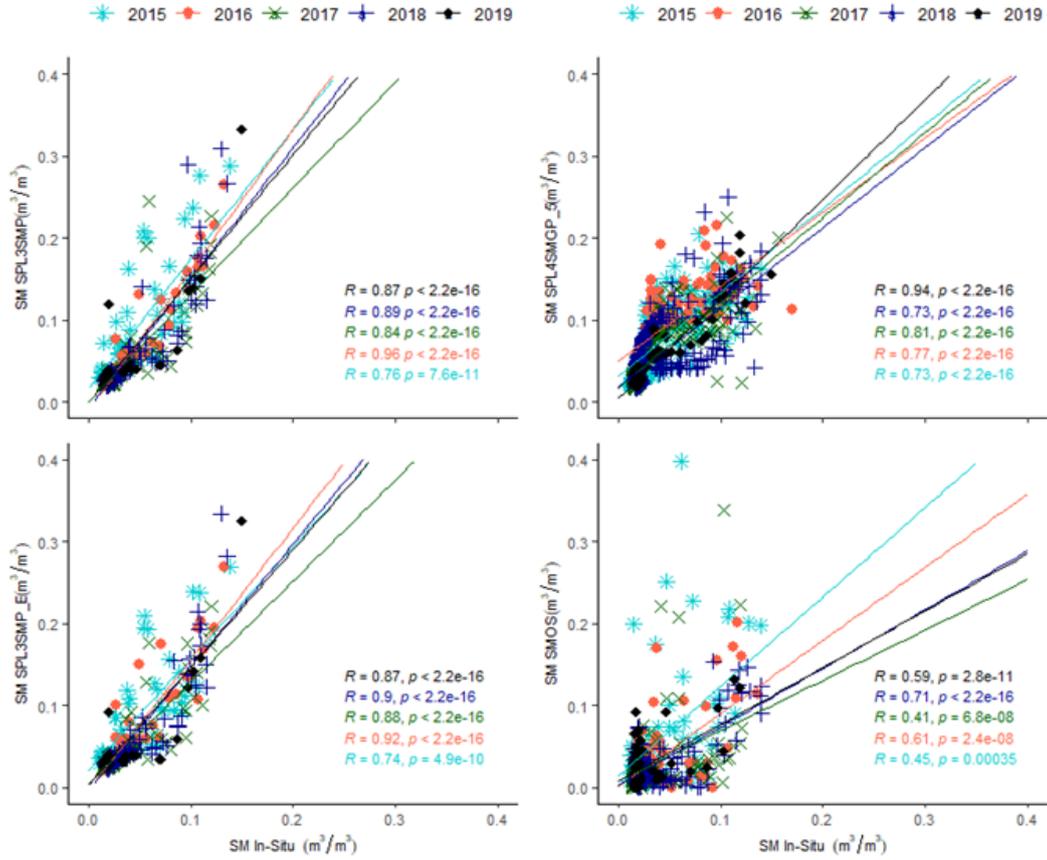


Figure 6. In situ SM versus SMAP and SMOS SM products in Bawdi site

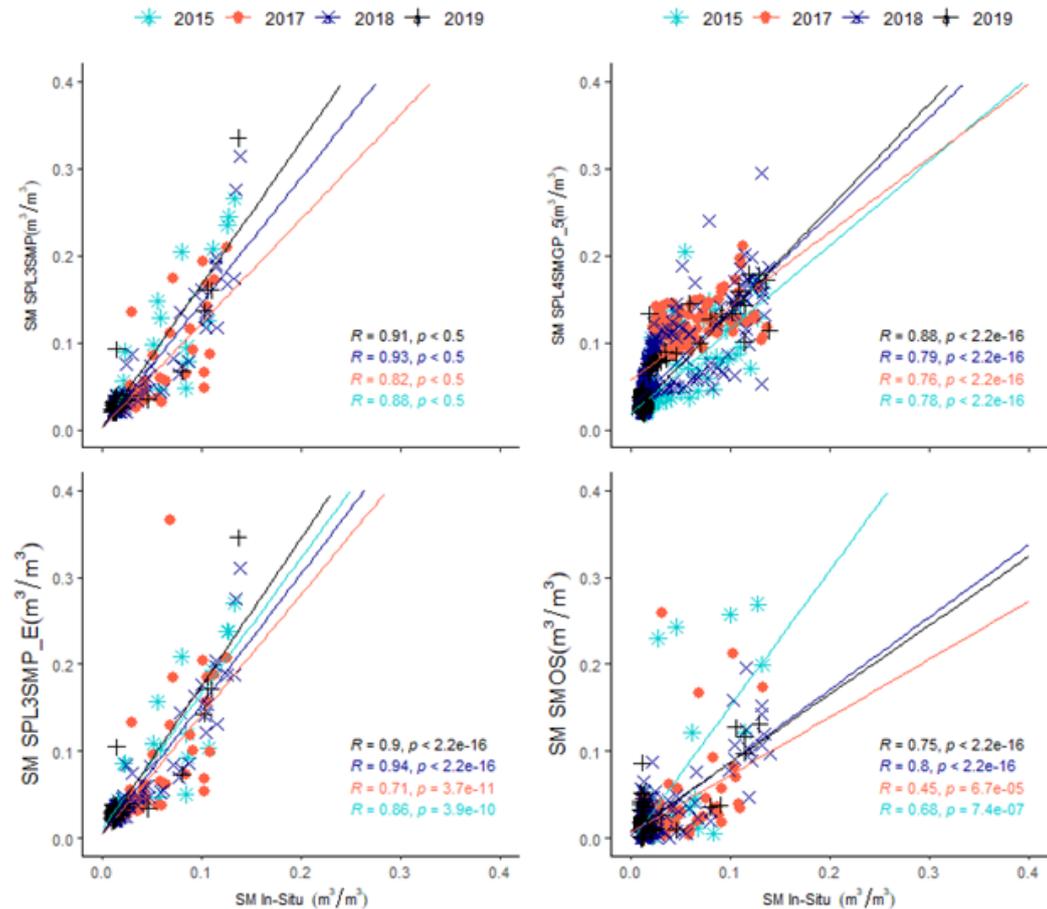


Figure 7. In situ SM versus SMAP and SMOS SM products Loumbi site

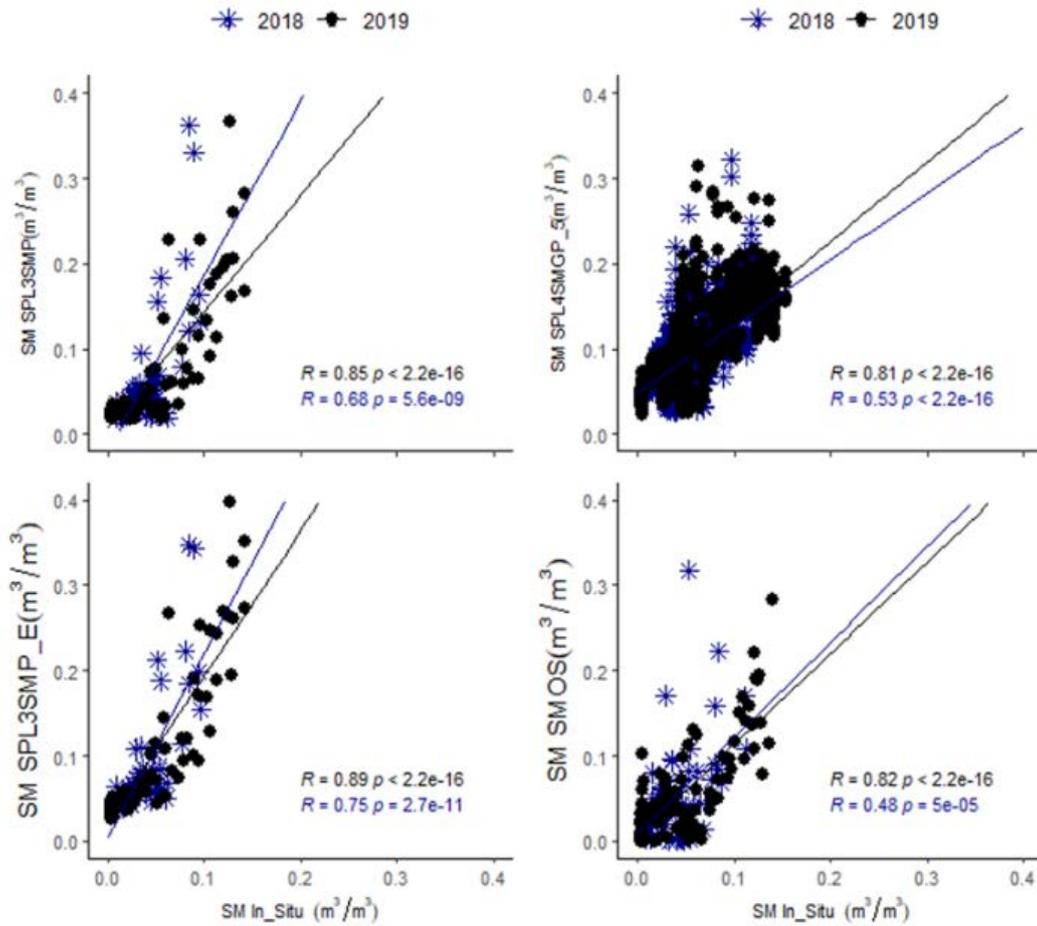


Figure 8. In situ versus SMAP and SMOS SM products at Niakhar site

• **Bawdi Site**

Figure 6 shows the scatter plot between satellite products and in situ SM for the period 2015-2019. We generally notice a good correlation between the SMAP products of the SPL3SMP level 3 sensor, version 7 with a spatial resolution of 36 km, and the in-situ SM with a correlation coefficient that varies over the five years between 0.76 and 0.96. With an RMSE smaller than 4% m<sup>3</sup>/m<sup>3</sup> for all years except 2015 (RMSE= 0.06 m<sup>3</sup>/m<sup>3</sup>) according to Table 4, confidence in the quality of the estimates of the different products is very high. Thus, despite its very low resolution of 36 km, the SPL3SMP products can reproduce in-situ SM at this site.

Furthermore, the annual correlation between the 9 km resolution product (SPL3SMP\_E) of SMAP and the in-situ SM is high and ranges from, 0.74 to 0.92. Except in 2015 when the RMSD is 0.05 m<sup>3</sup>/m<sup>3</sup>, the other years have an RMSE lower than 0.04 m<sup>3</sup>/m<sup>3</sup>. This shows the good performance of the SMAP sensor in estimating SM values.

For the SPL4SMGP\_5 9 km SMAP and SMOS sensors, the correlation coefficients range between 0.73-0.94 and 0.45-0.71, respectively (Table 4). The Bias values are also low for all sensors over the five years of study.

• **Loumbi Site**

In Loumbi (Figure 7), the study period is four years (2015, 2017, 2018, and 2019). Due to technical problems, we did not have data from 2016. The scatterplots for the Loumbi site are quite similar to those for Bawdi. Table 4 shows the statistical results between the microwave sensor

products and the ground measurements.

For the SMAP SPL3SMP sensor, we observe a good correlation with a correlation coefficient R between 0.86 and 0.93. The same remarks can be made as for the SPL3SMP\_E sensor of SMAP 9 km where the correlation coefficient varies between 0.86 and 0.94. We also find RMSE below 0.04 m<sup>3</sup>/m<sup>3</sup>. In addition, the SPL4SMGP sensor products have a very significant correlation while compared to the in-situ measurements with a correlation coefficient greater than 0.5 for all four years of study. However, the SMOS products are well correlated with the ground values, correlation coefficients above 0.5. The results obtained at the Loumbi site showed that the performance of the satellite sensors is very high for estimating SM despite RMSE sometimes exceeding the specifications (0.04 m<sup>3</sup>/m<sup>3</sup>).

• **Niakhar Site**

Figure 8 shows the scatter plots and Table 4 gives the RMSE, R, MAE, and Bias between all products. It is generally noted that the correlation coefficient varies between 0.48 and 0.89. A good correlation was noted between the SMAP SPL3SMP product and the ground observations between 2018 and 2019. It is also observed that a good part of the measurements is between 0.05 and 0.1 m<sup>3</sup>/m<sup>3</sup>. The largest correlation coefficient is found between the SMAP SPL3SMP\_E product and the ground values where R is equal to 0.89. The results found at the Niakhar site (agricultural area) show that there is a good link between the earth observation data and the in-situ data.

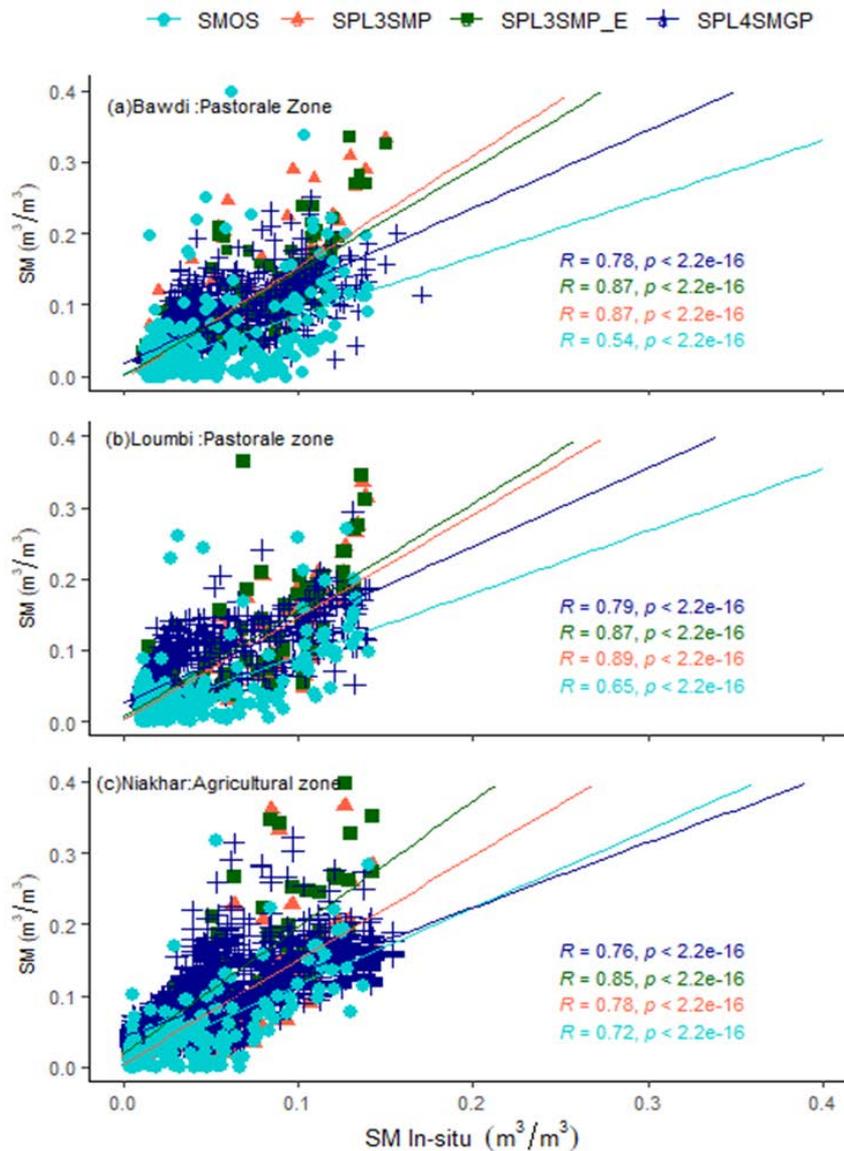
**Table 4.** Descriptive statistics between observations from SPL3SMP (36km), SPL3SMP\_E (9km), SPL4SMGP (9km), SMOS, and in-situ measurements from Bawdi, Loumbi, and Niakhar sites for a depth of 5cm: R, RMSE (in m<sup>3</sup>/m<sup>3</sup>), MAE (in m<sup>3</sup>/m<sup>3</sup>) and Bias (in m<sup>3</sup>/m<sup>3</sup>)

| Sites   | Products | SPL3SMP 36 km |      |      |      |     | SPL3SMP_E 9 km |      |      |      |     | SPL4SMGP_5 9 km |      |      |      |     | SMOS 0.25°x 0.25° |      |      |         |     |
|---------|----------|---------------|------|------|------|-----|----------------|------|------|------|-----|-----------------|------|------|------|-----|-------------------|------|------|---------|-----|
|         |          | R             | RMSE | MAE  | Bias | N   | R              | RMSE | MAE  | Bias | N   | R               | RMSE | MAE  | Bias | N   | R                 | RMSE | MAE  | Bias    | N   |
| Bawdi   | 2015     | <b>0.76*</b>  | 0.06 | 0.04 | 0.04 | 51  | <b>0.74*</b>   | 0.05 | 0.04 | 0.04 | 51  | <b>0.73*</b>    | 0.04 | 0.03 | 0.03 | 137 | 0.45              | 0.07 | 0.04 | 0.01    | 59  |
|         | 2016     | <b>0.96**</b> | 0.03 | 0.02 | 0.02 | 59  | <b>0.92**</b>  | 0.03 | 0.02 | 0.02 | 59  | <b>0.77*</b>    | 0.05 | 0.04 | 0.04 | 162 | <b>0.61*</b>      | 0.04 | 0.03 | -0.003  | 71  |
|         | 2017     | <b>0.84**</b> | 0.02 | 0.01 | 0.01 | 137 | <b>0.88**</b>  | 0.02 | 0.01 | 0.01 | 132 | <b>0.81**</b>   | 0.03 | 0.02 | 0.02 | 365 | 0.41              | 0.04 | 0.02 | -0.005  | 162 |
|         | 2018     | <b>0.89**</b> | 0.03 | 0.01 | 0.01 | 137 | <b>0.90**</b>  | 0.03 | 0.01 | 0.01 | 136 | <b>0.73*</b>    | 0.03 | 0.02 | 0.01 | 362 | <b>0.71*</b>      | 0.02 | 0.02 | -0.008  | 161 |
|         | 2019     | <b>0.87**</b> | 0.02 | 0.01 | 0.01 | 75  | <b>0.87**</b>  | 0.03 | 0.01 | 0.01 | 74  | <b>0.94**</b>   | 0.01 | 0.01 | 0.01 | 237 | <b>0.59*</b>      | 0.02 | 0.01 | -0.0004 | 105 |
| Loumbi  | 2015     | <b>0.88**</b> | 0.05 | 0.04 | 0.03 | 32  | <b>0.90**</b>  | 0.06 | 0.04 | 0.04 | 32  | <b>0.78*</b>    | 0.03 | 0.02 | 0.01 | 83  | <b>0.68*</b>      | 0.08 | 0.03 | 0.02    | 43  |
|         | 2017     | <b>0.82**</b> | 0.05 | 0.02 | 0.01 | 66  | <b>0.71*</b>   | 0.05 | 0.02 | 0.02 | 65  | <b>0.76*</b>    | 0.05 | 0.05 | 0.04 | 176 | 0.45              | 0.07 | 0.03 | -0.0003 | 73  |
|         | 2018     | <b>0.93**</b> | 0.03 | 0.01 | 0.01 | 137 | <b>0.94**</b>  | 0.03 | 0.02 | 0.02 | 136 | <b>0.79*</b>    | 0.04 | 0.03 | 0.03 | 365 | <b>0.80**</b>     | 0.02 | 0.01 | -0.0005 | 159 |
|         | 2019     | <b>0.91**</b> | 0.03 | 0.01 | 0.01 | 75  | <b>0.90**</b>  | 0.03 | 0.02 | 0.02 | 74  | <b>0.88**</b>   | 0.02 | 0.02 | 0.02 | 237 | <b>0.75*</b>      | 0.01 | 0.01 | 0.004   | 105 |
| Niakhar | 2018     | <b>0.68*</b>  | 0.06 | 0.03 | 0.02 | 57  | <b>0.75*</b>   | 0.07 | 0.05 | 0.05 | 57  | <b>0.53*</b>    | 0.05 | 0.04 | 0.04 | 153 | 0.48              | 0.05 | 0.03 | 0.01    | 64  |
|         | 2019     | <b>0.85**</b> | 0.04 | 0.02 | 0.02 | 123 | <b>0.89**</b>  | 0.06 | 0.04 | 0.04 | 122 | <b>0.82**</b>   | 0.05 | 0.04 | 0.03 | 365 | <b>0.82**</b>     | 0.03 | 0.02 | 0.01    | 155 |

### 3.3. Global Estimates of all Satellite Products versus in Situ Measurements

A site-by-site and global study between the measurements of the products from the satellite sensors and the ground observations were made and the results are shown in

Figure 9. Figures 9a, 9b, and 9c show the Bawdi, Loumbi, and Niakhar sites respectively. It is generally noted that many measurements are between 0.01 and 0.15 m<sup>3</sup>/m<sup>3</sup>. By grouping, all the measurements from all the sensors, a good correlation between satellite observations and field measurements can be observed at all sites.



**Figure 9.** Global interpretation between satellite products and in situ measurements: (a) Bawdi, (b) Loumbi, and (c) Niakhar

**Table 5. Soil moisture scores of all sites between satellite measurements versus in situ measurements**

| Soil moisture scores | SPL3SMP     |      |      | SPL3SMP_E   |      |      | SPL4SMGP_5  |      |      | SMOS |      |        |
|----------------------|-------------|------|------|-------------|------|------|-------------|------|------|------|------|--------|
|                      | R           | RMSE | Bias | R           | RMSE | Bias | R           | RMSE | Bias | R    | RMSE | Bias   |
| (a) Bawdi            | <b>0.87</b> | 0.03 | 0.02 | <b>0.87</b> | 0.03 | 0.02 | <b>0.78</b> | 0.03 | 0.02 | 0.54 | 0.04 | -0.002 |
| (b) Loumbi           | <b>0.89</b> | 0.04 | 0.02 | <b>0.87</b> | 0.04 | 0.02 | <b>0.79</b> | 0.04 | 0.03 | 0.65 | 0.04 | 0.03   |
| (c) Niakhar          | <b>0.78</b> | 0.05 | 0.02 | <b>0.85</b> | 0.06 | 0.05 | <b>0.77</b> | 0.05 | 0.04 | 0.72 | 0.03 | 0.01   |

Table 5 gives the R, RMSE, and Bias between satellite and in situ measurements. This may mean that soil and climatic characteristics do not have a great influence on the recovery of soil moisture measurements. With the SMAP products, a good correlation was respectively found at Bawdi where the correlation coefficient R is between 0.78 and 0.87 and RMSEs below the specifications ( $0.04 \text{ m}^3/\text{m}^3$ ); for SMOS, R is equal to 0.54, RMSEs  $< 0.04 \text{ m}^3/\text{m}^3$ . At Loumbi where R is between 0.79 and 0.89 for SMAP measurements and 0.65 for SMOS measurements. These two sites together give Ferlo which has similar scores to these two different sites,  $0.77 < R < 0.89$ .

In Niakhar, very good results were also obtained with all satellite sensor products as the correlation coefficient is between 0.77 and 0.85 for SMAP and 0.72 for the SMOS product. The results obtained in the Ferlo (silvopastoral zone) corroborate the results of [45].

## 4. Validation of Soil Moisture in the Root Zone

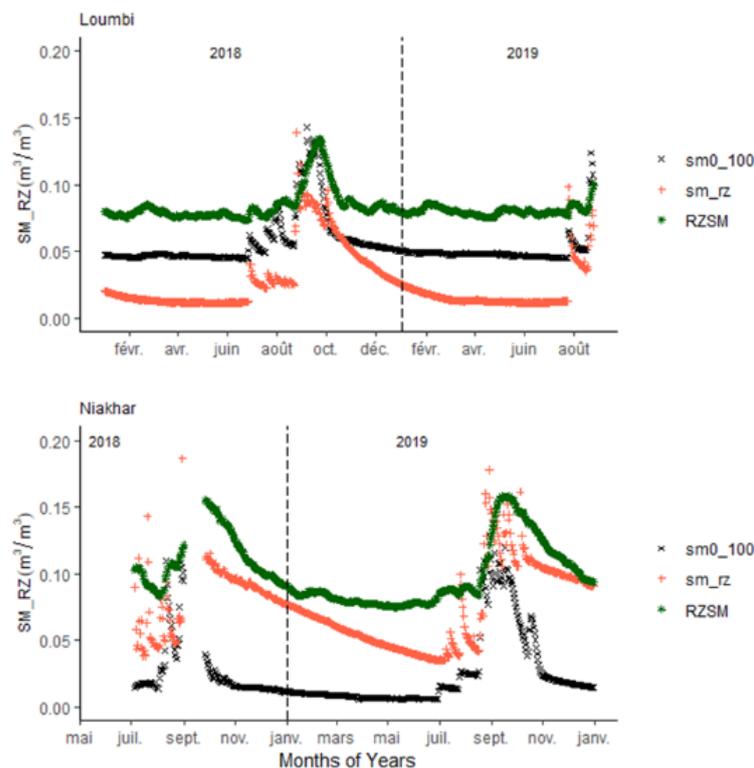
### 4.1. Loumbi: Pastoral Area

Figure 10 shows the temporal dynamics of soil moisture from 2018 to 2019. Integrated moisture with depths of 5, 10, 40, and 100 cm is calculated to have an in-situ root

zone soil moisture. This calculated moisture is compared to the moisture of the SMAP and SMOS sensors in the root zone. Generally, all observations reflect the seasonal dynamics, i.e., the two seasons can be distinguished. However, there are also differences between the SMAP values and the other types of measurements, especially during the dry season. The SMOS curve is above all other curves; this could indicate that this sensor overestimates the soil moisture measurements of the root zone.

### 4.2. Niakhar: Agricultural Areas

Figure 10 also shows the temporal soil moisture dynamics of the root zone of the satellite products and the in-situ measurements at Niakhar. This site is in an agricultural environment with soil that is constantly plowed. There is a difference in the variation of moisture values from the microwave sensors compared to the Loumbi site. Here, the difference in values is noted between the satellite and in-situ products especially during the dry season when the microwave observations are significantly higher. In September 2018, in the middle of the rainy season, the soil moisture variations of the SMAP sensor are much higher than the other measurements with a peak above  $0.3 \text{ m}^3/\text{m}^3$ . This can be justified by the intensity of the rainfall and the capacity of the soil to infiltrate. In agricultural areas, water infiltration is favored by anthropogenic activities in the soil.



**Figure 10.** Spatial-temporal variation of SM of the root zone in a silvopastoral area (Loumbi) and an agricultural area (Niakhar)

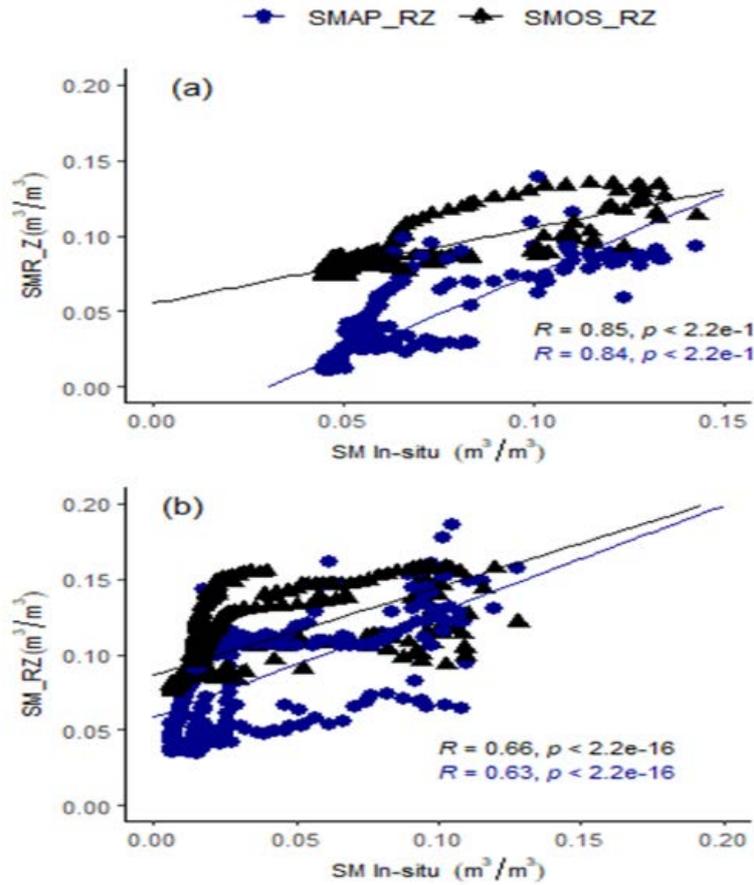


Figure 11. Comparison between microwave sensor products and in situ root zone soil moisture measurements: (a) Loumbi and (b) Niakhar

Table 6. Correlation coefficient, RMSD, MAE, and Bias between the root zone soil moisture of the SMAP\_RZ resolution 9 km, SMOS\_RZ sensors, and the in-situ soil moisture in Loumbi and Niakhar

| SM scores | SMAP_RZ     |      |       | SMOS_RZ     |      |      |
|-----------|-------------|------|-------|-------------|------|------|
|           | R           | RMSE | Bias  | R           | RMSE | Bias |
| Loumbi    | <b>0.85</b> | 0.03 | -0.03 | <b>0.84</b> | 0.03 | 0.03 |
| Niakhar   | <b>0.63</b> | 0.05 | 0.05  | <b>0.66</b> | 0.08 | 0.07 |

Figure 11(b) shows the scatter plots and Table 6 the scores between the different products. A good correlation was found where the correlation coefficient is equal to 0.84 for SMAP\_RZ and 0.85 for SMOS\_RZ. This demonstrates the good performance of SMAP and SMOS in taking SM measurements of the root zone.

### 5. Conclusion

In this study, a multi-year evaluation of various SM microwave products of the surface (5 cm) and the root zone (0-100 cm) soil layers was carried out over three sites located in Senegal. Overall, the results obtained in this study corroborate the results found by previous studies in the Sahelian zone [15,16,76].

Microwave sensors that scan the first few centimeters of soil depth (5cm) have products that have mostly good correlation and RMSEs that are often below specifications, i.e., 0.04 m3/m3 in a pastoral area: Loumbi and Bawdi. Therefore, the products of SMAP and SMOS sensors, with high correlation coefficients in these two sites, can be good indicators of surface soil moisture. Similarly in the

agricultural site of Niakhar, satellite results and situ surface measurements show the sensitivity of the sensors in the agricultural area.

Due to lack of data at the Bawdi site, root zone estimates were made at Loumbi and Niakhar with the SPL4SMGP 9 km sensor from SMAP\_RZ and the SMOS\_RZ sensor from SMOS. A very large result was found at Loumbi with RMSEs mainly less than or equal to 0.04 m3/m3. The results found at the Niakhar site also show the importance of microwave sensors on the recovery of soil moisture measurements from the root zone. Thus, this work shows the sensitivity of in situ and satellite sensors on L-band soil moisture recovery.

Nevertheless, further studies are needed to improve the observations of microwave sensors on soil moisture retrieval for better agricultural monitoring in the Sahel, particularly in Senegal.

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